

A HIGH RESOLUTION GLOBAL OCEAN MODEL with VARIABLE FORCING of WIND, HEAT, and FRESHWATER: II) TRANSPORT VARIABILITY

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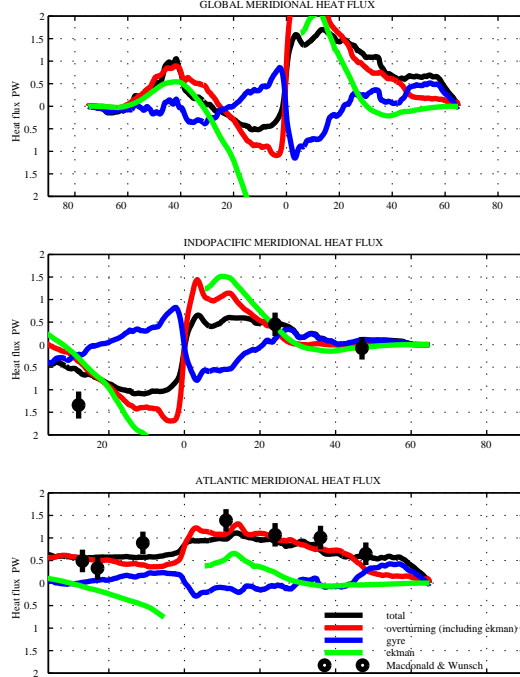
Goal: Understand ocean's low freq. variability

Conclusions in Red

Model Simulation Details:

- (Semtner & Chervin '92, Stammer, et al. '96, Tokmakian, '96)
- 1/4° avg. Semtner/Chervin Primitive Eq. OGCM
- Parallel Ocean Climate Model POCM-vers 4C
- Forced with ECMWF reanalysis + oper. 1979-1996
- > heat, freshwater, wind stress - varying daily (including rivers)

Meridional Heat Flux



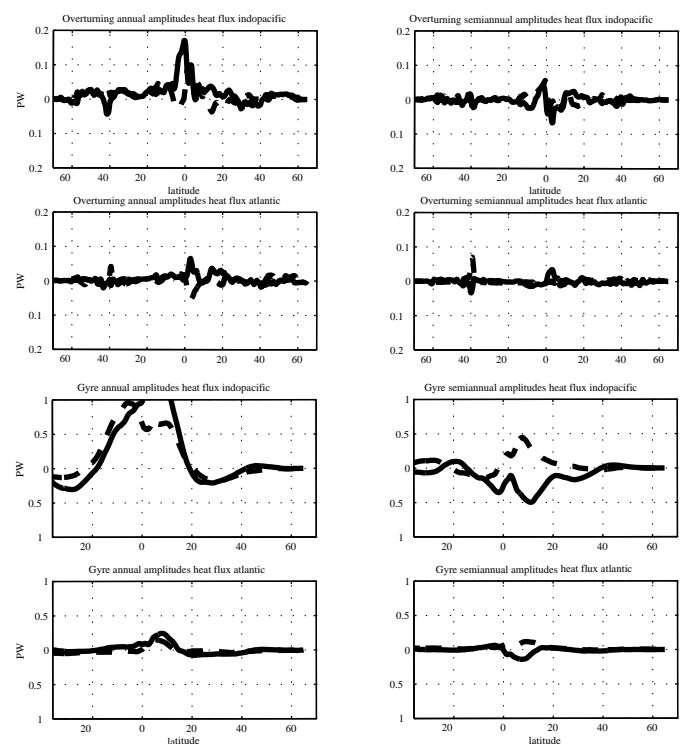
Calculated from monthly means
Overturning, where $\langle \rangle$ is mean
$$\rho^* C_p \int_H [L \langle V_{zonal} \rangle \langle T_{zonal} \rangle] dz,$$

Gyre
$$\rho^* C_p \int_H (V - \langle V \rangle) (T - \langle T \rangle) d\lambda dz$$

Ekman
$$C_p \int_L -\tau' / f (T_{surface} - \langle T_{zonal} \rangle) d\lambda$$

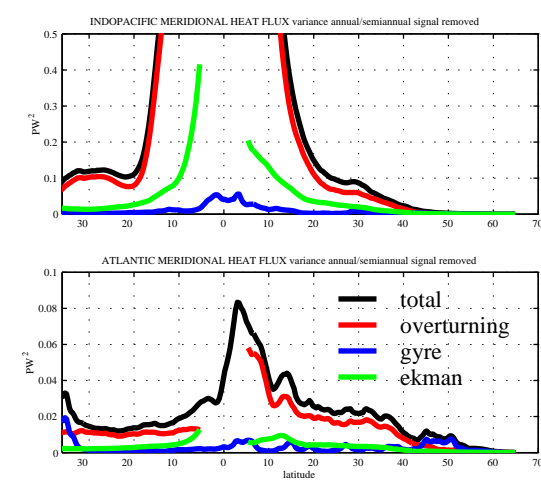
Long Term Mean

POCM 4C simulates the mean heat flux well as compared to Macdonald & Wunsch ('96).



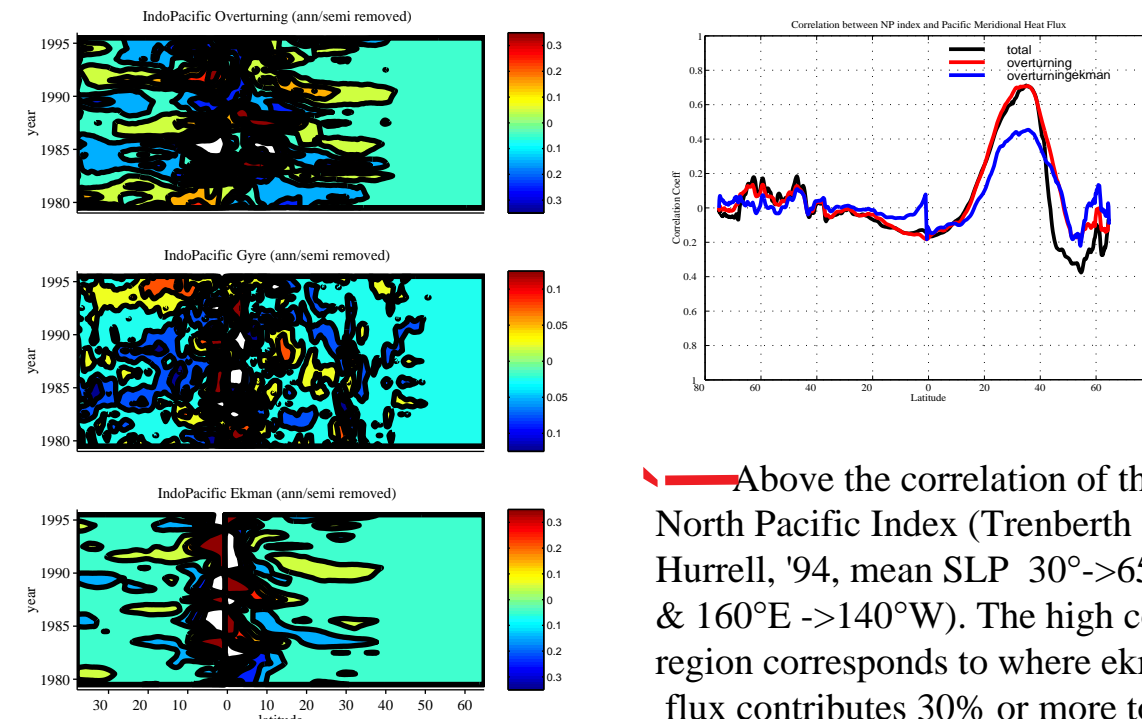
The plots to the left show the amplitudes of the annual and semi-annual signals for the components of the heat flux in the Indo-Pacific and the Atlantic regions.

The annual cycle is 8% of the total HF at 24°N in the overturning component; 30% of total for the gyre component at 24°N; semi-annual contributes less than 1%.



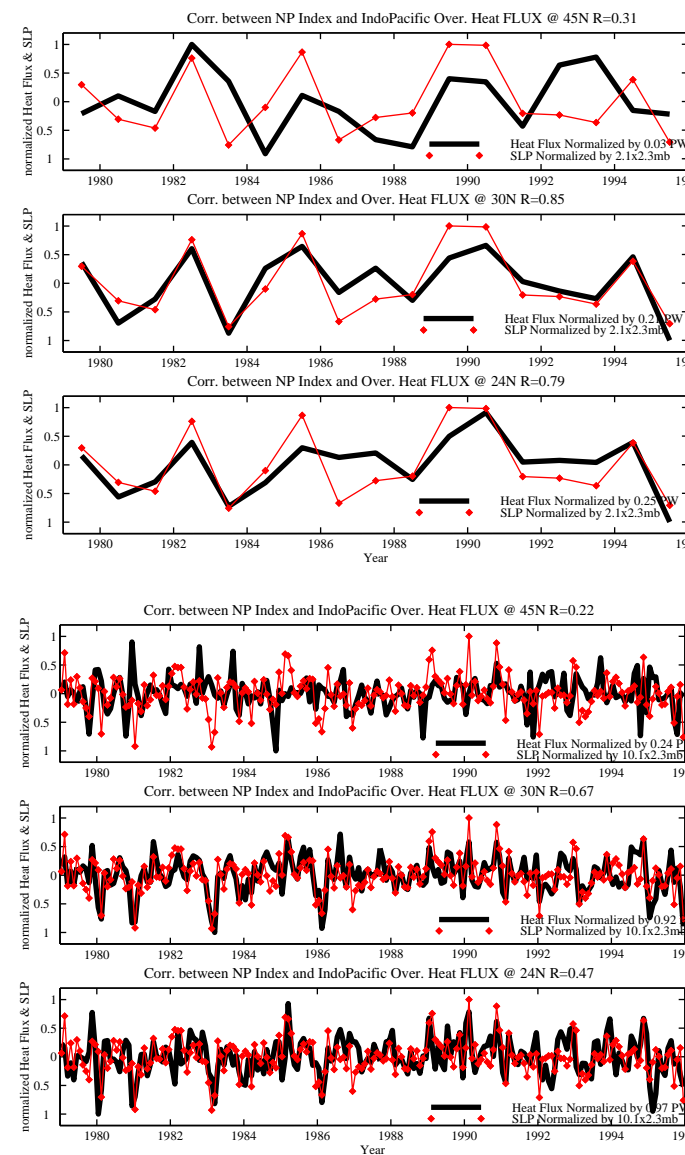
The diagram to the left shows the variance in the heat flux with the annual and semi-annual signal removed, for each component.

Indo-Pacific Meridional Heat Flux



Time-latitude plots are shown for the components of heat flux. Note, the annual and semi-annual signal has been removed and the ekman component is included in the overturning component. Low frequency variability > annual is clearly simulated by the model visible in these integrated quantities. (Contours: gyre 0.05 PW, other 0.1PW).

At 24°N, the top 400 m contribute most to the variability of the overturning.



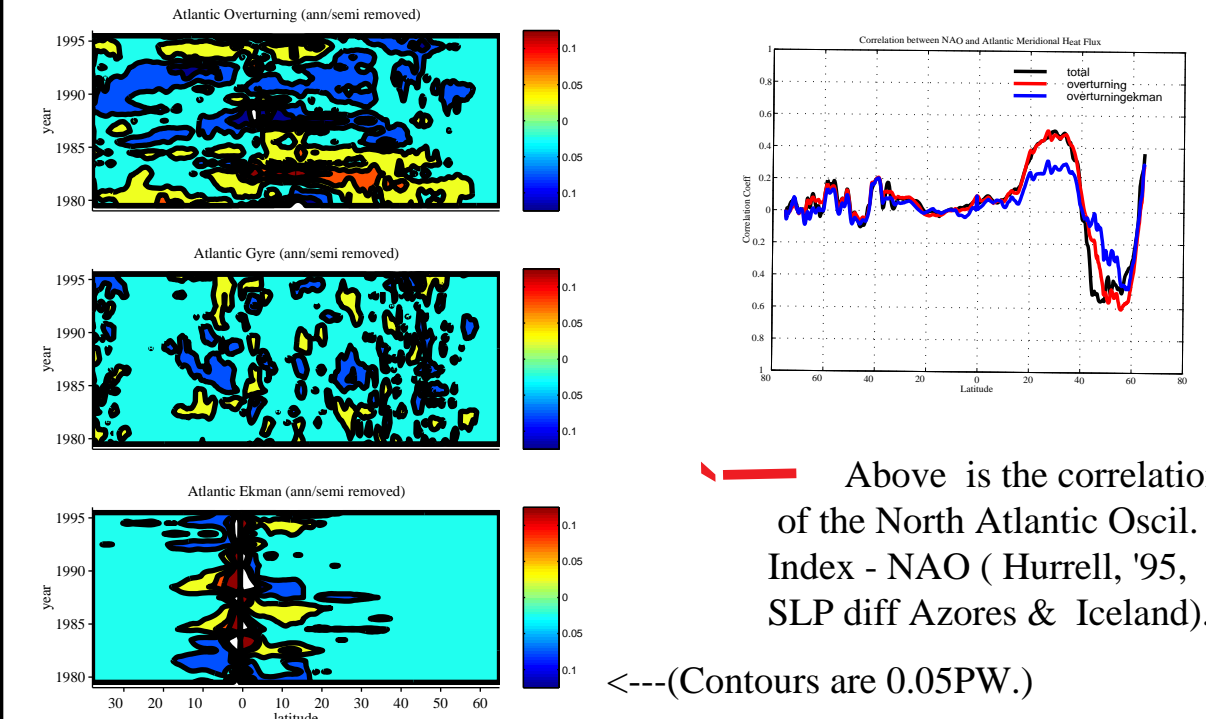
Shown at left are normalized time-series of the NP index (with annual/semi-annual signal removed) and the meridional heat flux at three latitudes, 45°, 30°, & 24°N.

Maximum variability is on the order of 0.4 PW at 24°N, & the monthly minimum is found in the during spring/summer.

@24°N; total HF in Pacific can be separated as:

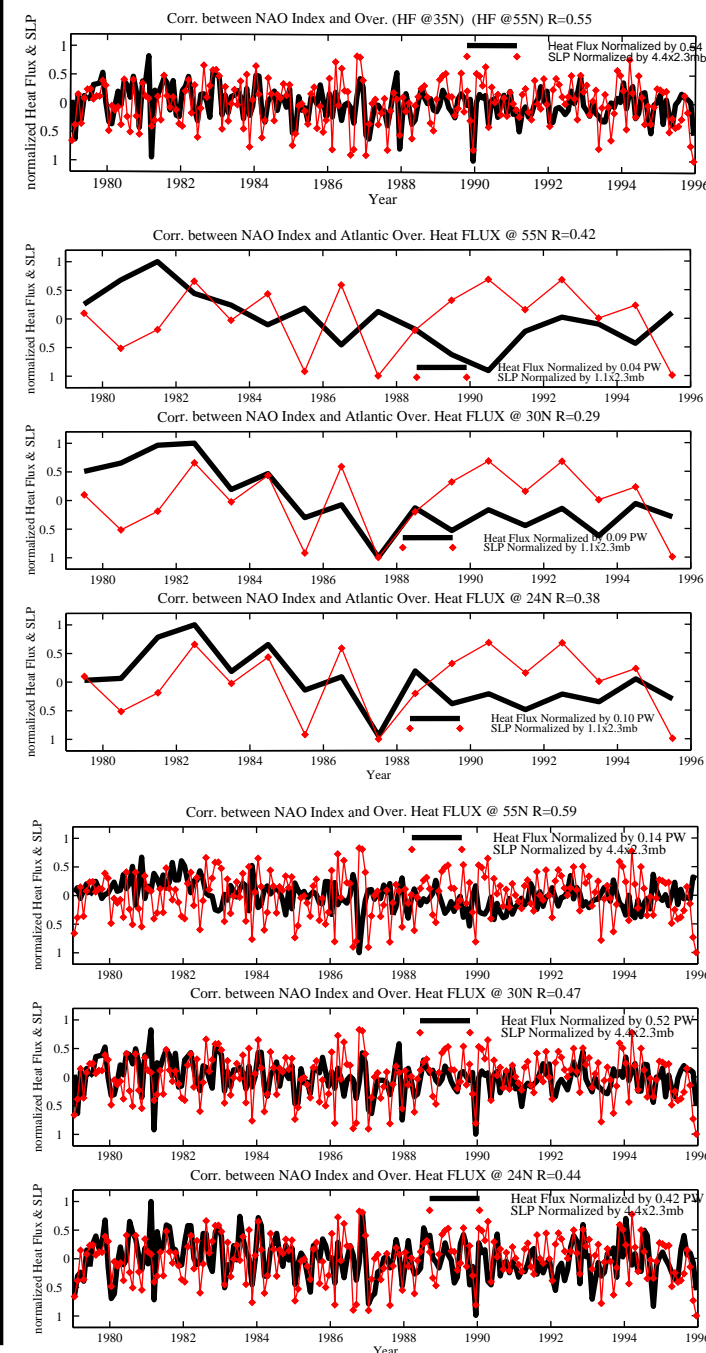
Ekman = 0.38 PW
Kuroshio (gyre) = -0.026PW
Kuroshio (over.) = 2.4PW
Mid-ocean (gyre) = -0.03PW
Mid-ocean (over) = -2.24PW
0.48 PW

Atlantic Meridional Heat Flux



Time-latitude plots are shown for the components of heat flux. Note, the annual and semi-annual signal has been removed and the ekman component is included in the overturning component.

The ekman flux contributes little to the overturning heat flux north of 40°N.

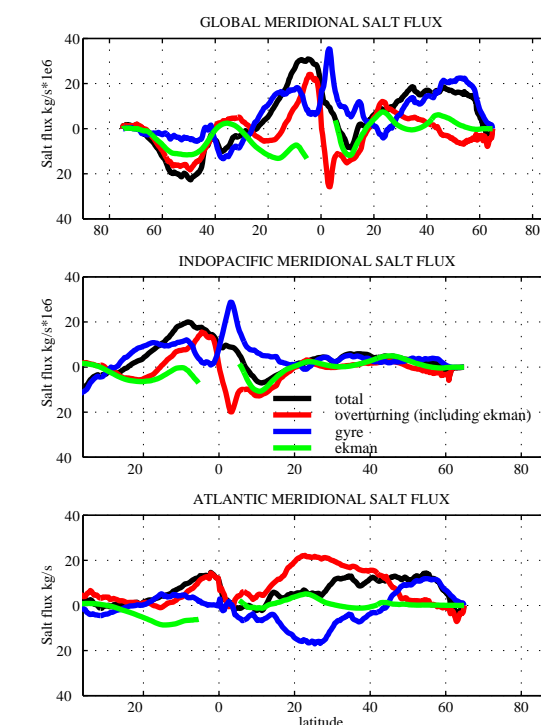


A time-series of the NAO index (with annual/semi-annual signal removed) and the difference in the meridional heat flux at 35°N and 55°N (because NAO is a difference in SLP). This shows the relative influence of atmospheric conditions on the overturning heat flux of the model.

To the left are normalized time-series of the NAO index (with annual/semi-annual signal removed) and the meridional heat flux at three latitudes, 55°, 30°, & 24°N.

Maximum annual variability is on the order of 0.2 PW at 24°N, & the monthly variability also 0.2 PW.

Meridional Salt & Freshwater Fluxes



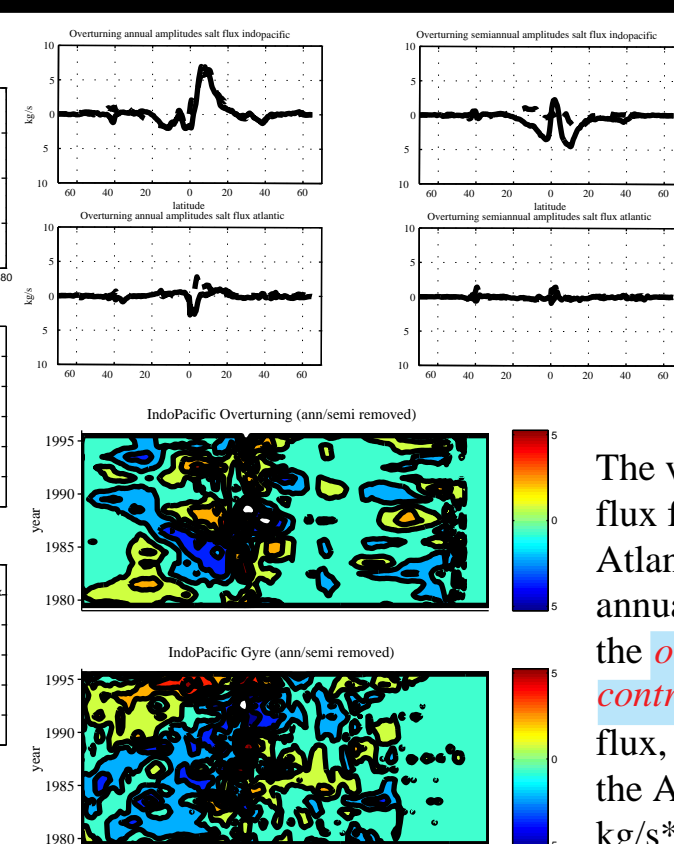
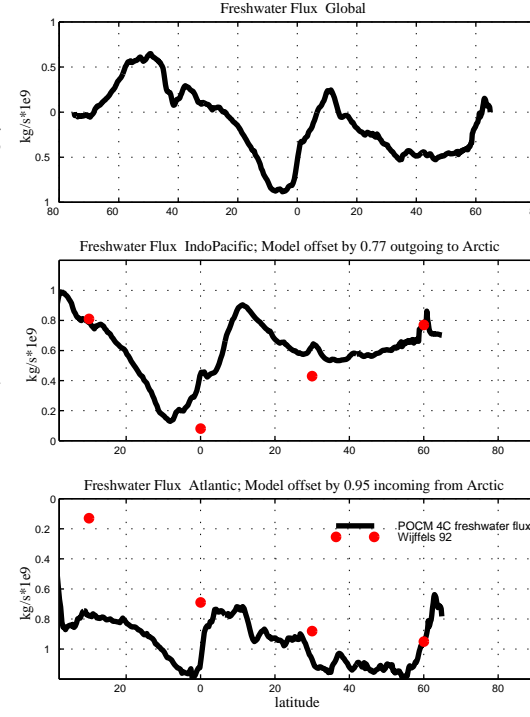
The salt flux calc. follow the heat flux calc. above, replacing T with S, & changing the constant to ρ only.

Freshwater Calc. (mass flux minus % which is salty)
$$\rho (\langle v \rangle - \langle v \rangle \langle S \rangle - \langle v' S' \rangle)$$

(Arctic inflow/outflow included at 60°N.)

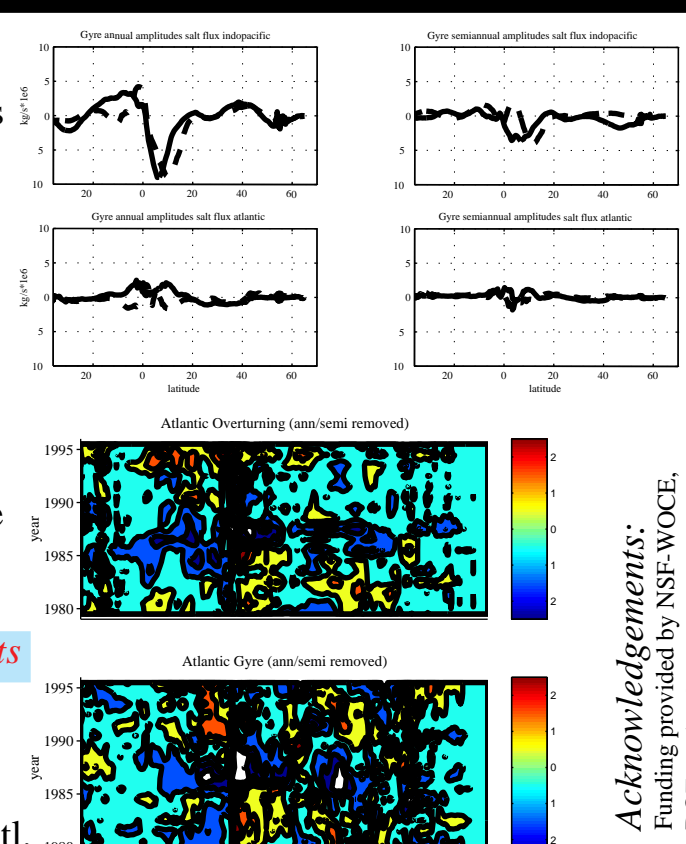
Plot to right shows the freshwater trans. in the model as compared to Wijffels '92. The model does not simulate the northward flux of Antarctic water correctly

N. Atlantic & Indo-Pacific fluxes are similar to observations



The plots to the right & left show the amplitudes of the annual and semi-annual signals for the components of the salt flux in the Indo-Pacific and the Atlantic regions.

The variability with time of the salt flux for the Indo-Pacific and the Atlantic with the annual/semi-annual signal removed shows that the overturning and gyre components contribute equally to the total salt flux, especially in the S. Pacific and the Atlantic. Contours are 1.5 kg/s*1e6: Indo-Pac; kg/s*1.e6 for Atl.



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